



Seismic magnitudes, entropy and b -value

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Received: 10 May 2024 / Accepted: 16 July 2024

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Abstract

A closed relationship between the Gutenberg–Richter b -value (or $\beta = b \ln 10$) and the information or Shannon entropy is found and checked through numerical evaluation of the entropy using exact probabilities derived directly from the magnitude exponential distribution. Comparison of the numerical evaluation of the entropy over a finite magnitude range makes it possible to assess the possible contribution to the entropy of real or hypothetical very large magnitudes, and these contributions are found to be quite small. The relationship is also compared with entropies calculated from synthetic data, and Monte Carlo simulations are used to explore the behavior of entropy determinations as a function of sample size. Finally, it is considered how, for the usual case of having data from a single realization, i.e., a single magnitude data set, since estimates of the entropy and of the Aki–Utsu b -value are measured in different ways, they are not redundant and may be complementary and useful in determining when a sample is large enough to give reliable results.

Keywords Information entropy · Gutenberg–Richter b -value · Seismic magnitudes

Introduction

The Shannon, or information, entropy and the b -value of the Gutenberg–Richter distribution, both discussed in detail below, have become useful and widely used tools in the study of seismicity, because both seem to quantify behaviors of seismicity related to the levels of stress. Here, a relationship between b and the entropy of the seismic magnitudes will be presented, some of its features will be discussed, and ways in which these measures can complement each other will be proposed.

In what follows unrounded magnitudes will be denoted by m and magnitudes rounded to ΔM by M (usually $\Delta M = 0.1$).

The G–R b -value

Ishimoto and Ida (1939) and Gutenberg and Richter (1944) observed that seismic magnitudes are distributed as:

$$\log_{10} N(M) = a - b(M - M_c); M \geq M_c, \quad (1)$$

where $N(M)$ is the number of magnitudes $\geq M$, and b describes the proportion of large magnitudes to small ones (Richter 1958). The magnitude origin has been shifted by M_c , the completeness magnitude, below which $\log_{10} N(M)$ ceases to behave linearly due to insufficient coverage (e.g., Wiemer and Wyss, 2000). Although the physical meaning of the Gutenberg–Richter relation, and of related distributions of seismic energy and moment are still subject to discussion (e.g., Wyss 1973; El-Isa and Eaton 2014), the b -value has been widely used to characterize seismicity in different regions in the world (e. g. Kagan 1999; Utsu 2002), and it has been proposed that b is related to the fractal dimension (Aki 1981; Hirata 1989; Wyss et al. 2004). There are many studies that relate b inversely to the level of stress and observe decreases in its value before large earthquakes (DeSalvio and Rudolph 2021; El-Isa and Eaton 2014; Enescu and Ito 2001; Frohlich and Davis 1993; Godano et al. 2024; Hu et al 2024; Lacidogna et al., 2023; Li and Chen 2021; Nanjo et al. 2012; Scholtz 2015; Schorlemmer et al 2005; Utsu 2002; Wang 2016; Wyss 1973; Wyss et al. 2004; and many others), which gives b an important role in earthquake hazard estimation and forecasting.

b -values can be estimated directly from the slope of the linear range on the G–R histogram (e.g., Guttorp 1987;

Edited by Prof. Ramón Zúñiga (CO-EDITOR-IN-CHIEF).

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Monterroso and Kulhanek, 2003), but frequently b -values are estimated from the mean magnitude (Aki 1965; Marzocchi and Sandri 2003; Tinti and Mulargia 1987; Utsu 1965), and most studies use the Aki–Utsu maximum likelihood estimate:

$$b = \frac{\log_{10}(e)}{\bar{M} - m_c}, \quad (2)$$

where \bar{M} is the observed mean of the data (Aki 1965; Utsu 1965). This estimate shares with the entropy determinations the problem of determining m_c , but otherwise it is based on the mean magnitude that, in a way, incorporates the information from all magnitudes. This measure is extremely easy to obtain but, unfortunately, many people use (2) as a magic formula without considering that the estimate will be good only if the observed \bar{M} is close to the mean of the distribution μ (compare (2) with (7)), which requires having a sample large enough to be representative (Geffers et al. 2022; Marzocchi et al. 2020; Nava et al. 2017; Ogata and Yamashina 1986; Shi and Bolt 1982).

The information entropy

Another important statistical–probabilistical concept is Shannon’s definition of the information entropy, S , of a process characterized by K states or classes of events, each having probability P_i , with:

$$\sum_{i=1}^K P_i = 1, \quad (3)$$

as:

$$S = - \sum_{i=1}^K P_i \log_2 P_i \equiv \sum_{i=1}^K s_i \quad (4)$$

(Shannon 1948), where the logarithm can have any base; we will use base 2 because it is the one most commonly used for information purposes and yields an entropy expressed in bits, easy to interpret. Capital letters have been used for the probabilities to emphasize that they are not densities, and in this definition it is implicitly assumed that $0 \leq P_i \leq 1$, so that $\log_2 P_i \leq 0 \forall i$. Each term in the first sum in (4) is the contribution to the total entropy S of the probability of each rounded magnitude class, called entropy score by Harte and Vere-Jones (2005), and will be denoted by s_i , where i is the index of the class, or generally as s .

In Fano (1961) the self-information of an event with probability P_i is given by:

$$I_i = - \log_2 P_i, \quad (5)$$

the entropy (4) can be recognized as the expected self-information of the process. Although the self-information ranges from zero to infinity, the contribution to the entropy from any probability ranges from zero, for both $s(0)$ and $s(1)$, to the maximum $s(e^{-1}) = 0.530738$ bit, as shown in Fig. 1. This point will be retaken later.

The concept of entropy has been widely used in seismology, particularly through the Principle of Maximum Entropy (PME), to study distributions, recurrence relationships, model stress fields, estimate seismic hazard, etc. (Bookstein 2021; Berrill and Davis 1980; De Santis et al. 2011; Dong et al. 1984; Feng and Luo 2009; Janes, 1957; Main and Naylor 2008; Mansinha and Shen 1987; Shen and Mansinha 1983; Telesca et al. 2004). Other studies use entropy as an indicator of proximity to criticality (Main and Al-Kindy 2002; Vogel et al. 2020), some using so-called natural time (Ramírez-Rojas et al. 2018; Rundle et al. 2019; Sarlis et al. 2018; Varotsos et al. 2004; Varotsos et al., 2022; Varotsos et al. 2023), many using other definitions of entropy, and some for seismic electric signals (Varotsos et al. 2006). Entropy has also been used to study the spatial distribution of seismic sources (e.g., Bressan et al. 2017; Goltz, 1966; Goltz and Böse 2002; Nava et al. 2021; Nicholson et al. 2000; Ohsawa 2018) and to study noise (e.g., Lyubushin 2021).

It is because of the possible usefulness of both the b -parameter and the magnitude entropy that it is important to explore the relationship between these two observables.

The entropy of seismic magnitude distributions

Let the process considered in the information entropy be the seismic magnitudes and the classes be the classes of a magnitude histogram, and let us see what can the entropy be

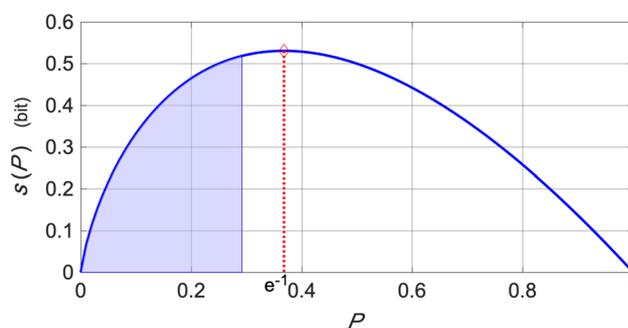


Fig. 1 Contribution of each particular probability value to the entropy. The dotted line indicates the position of the maximum for $P = e^{-1}$. The shaded area indicates the entropy for the range of probabilities corresponding to magnitude distributions with $b \leq 1.5$ and $\Delta M = 0.1$ (discussed below)

expected to be like by assuming that the magnitudes obey a G–R distribution.

The G–R relation (1) is a reverse cumulative histogram corresponding to an exponential magnitude probability density function,

$$p(m) = \beta e^{-\beta(m-m_c)}, \quad (6)$$

$$\beta = b \ln(10) = 1/(\mu - m_c), \quad (7)$$

where μ is the mean of the exponential distribution, and, since it should be related to all magnitudes that contribute to the rounded ones, is defined for unrounded magnitudes $m \geq m_c$, where $m_c = M_c - \Delta M/2$.

Let the classes considered in (3) correspond to the magnitudes rounded to $\Delta M = 0.1$, and let the probability of the class of a given rounded magnitude M_i , where $M_1 = M_c$, be P_i .

Commonly, P_i is approximated from (6) as:

$$P_i \approx p(M_i) \Delta M \quad (8)$$

(e.g., Rundle et al. 2019); a better procedure will be proposed below, but for now let us digress to discuss some reported results based on this approximation.

The Entropy of a Continuous Distribution

Substitution of (8) in (4) yields:

$$S = - \sum_{i=1}^K p(M_i) \Delta M \log_2 [p(M_i) \Delta M], \quad (9)$$

which can be written as:

$$S = - \sum_{i=1}^K p(M_i) \log_2 p(M_i) \Delta M - \log_2 \Delta M. \quad (10)$$

On letting $\Delta M \rightarrow 0$ the first term on the right side of (10) becomes what Shannon (1948) defined as the *entropy of a continuous distribution for a process having probability density distribution* $p(m)$:

$$S^c = - \int_{-\infty}^{\infty} p(m) \log_2 p(m) dm, \quad (11)$$

which we will denote by S^c to differentiate it from what would be the limit of the entropy in (10). Formula (11), without the minus sign, corresponds to what Wiener (1948) defined as *the amount of information of* $p(M)$, not as entropy. Shannon (1948) states that “The entropies of continuous distributions have most (but not all) of the properties of the discrete case.”, and it is clear they differ in this case, because the second term on the right-hand side of (10) has not been

included in the limit and this term grows as ΔM decreases and tends to infinity as $\Delta M \rightarrow 0$ (Mansinha and Shen 1987). Goldman (1953) is aware of the $-\log_2 \Delta M$ term, but states that it cancels out, which is certainly not the case for the problem at hand. Thus, S^c (11) is not the limit of S (4).

A problem with S^c is that the meaning of $-\log_2 p(m)$ is not clear, because the definitions of self-information and information entropy refer to mass probabilities, not to densities. For exponential distributions, unless $\beta < 1$, i. e., $b < 1/\ln 10 \approx 0.43429448$, which is an unrealistic value, the integral in (11) will include a range with $p(m) > 1$ that would imply negative information and contribute negative entropy.

Equation (11) has been used in several studies (e.g., De Santis et al. 2011; Main and Burton, 1964; Posadas et al., 2002; Posadas et al. 2021; Shen and Mansinha 1983) with varying results, some of them not applicable to the original definition of information entropy. For example, De Santis et al. (2011) obtained the relationship $S^c = \log(\text{eloge}) - \log b$, which, although appropriate for (11), differs from entropy estimates obtained from (4), implies an unrealistic upper limit $b_{max} = \text{e} \log_{10} e = 1.1805$, and so illustrates the perils of using (11). A useful relationship between entropy and b needs to be based on the original definition of entropy (4).

Entropy of exponential distributions and its relationship with b

Although this paper is oriented toward seismic magnitude distributions, what follows is applicable to any exponential probability distribution.

Coming back to Eq. 4, instead of using the approximation (8), the exact probability corresponding to the class of a rounded magnitude M_i can be calculated exactly as:

$$P_i \equiv P(M_i) = \int_{M_i - \Delta M/2}^{M_i + \Delta M/2} \beta e^{-\beta(m-m_c)} dm, \quad (12)$$

which results in:

$$\begin{aligned} P_i &= e^{-\beta(M_i - m_c)} \left(e^{\beta \frac{\Delta M}{2}} - e^{-\beta \frac{\Delta M}{2}} \right) \\ &= e^{-\beta(M_i - M_c)} (1 - e^{-\beta \Delta M}) \equiv e^{-\beta(M_i - M_c)} \Delta M p. \end{aligned} \quad (13)$$

To show how this probability estimation compares with the approximation shown before in 2.1, (8) can be written as:

$$\beta e^{-\beta(M_i - m_c)} \Delta M = e^{-\beta(M_i - M_c)} \beta e^{-\beta \Delta M/2} \Delta M, \quad (14)$$

so both (13) and (14) consist of the same exponential multiplied by different factors that are shown in Fig. 2 for various values of b . Both factors differ by very little for small

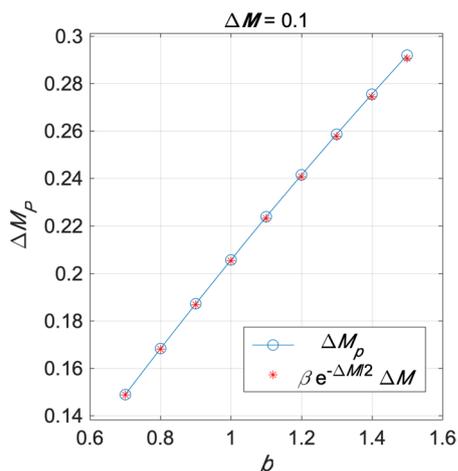


Fig. 2 Comparison of the factors that multiply an exponential to evaluate or estimate probabilities

b-values, but for large *b*-values ΔM_p is appreciably larger than the factor in (14), which shows that it is worthwhile to use the exact probability from (13).

Substituting probability (13) in (4) yields:

$$S = - \sum_{i=1}^K e^{-\beta(M_i - M_c)} \Delta M_p \log_2 \left[e^{-\beta(M_i - M_c)} \Delta M_p \right], \quad (15)$$

which is the expression for the entropy that will be used to calculate the theoretical entropy corresponding to a given magnitude distribution, to illustrate how the elements of the magnitude distribution contribute to the entropy, and to estimate through Monte Carlo simulation, what can be expected from data samples of different sizes.

To obtain an estimate for the theoretical value of *S*, let $K \rightarrow \infty$ in (15) because the theoretical G–R distribution does not have an upper limit; this limit will be discussed below. Equation (15) can then be written as:

$$S = - \sum_{i=1}^{\infty} e^{-\beta(M_i - M_c)} \Delta M_p \left[-\beta(M_i - M_c) \log_2 e + \log_2 \Delta M_p \right],$$

or:

$$S = \Delta M_p \log_2 e \sum_{i=1}^{\infty} -\beta(M_i - M_c) e^{-\beta(M_i - M_c)} - \log_2 \Delta M_p \sum_{i=1}^{\infty} e^{-\beta(M_i - M_c)} \Delta M_p. \quad (16)$$

The sum in the second right-hand term of (16) is the total probability equal to unity. In the first right-hand term, the factor $(M_i - M_c)$ takes values $0\Delta M, 1\Delta M, 2\Delta M, 3\Delta M, \dots$, so the sum written explicitly as:

$$0 - 1\Delta M\beta e^{-1\Delta M\beta} - 2\Delta M\beta e^{-2\Delta M\beta} - 3\Delta M\beta e^{-3\Delta M\beta} - \dots, \quad (17)$$

can be recognized as the series representation of:

$$\Delta M\beta \frac{d}{dx} (1 - e^x)^{-1} = \Delta M\beta \frac{e^x}{(1 - e^x)^2}, \quad (18)$$

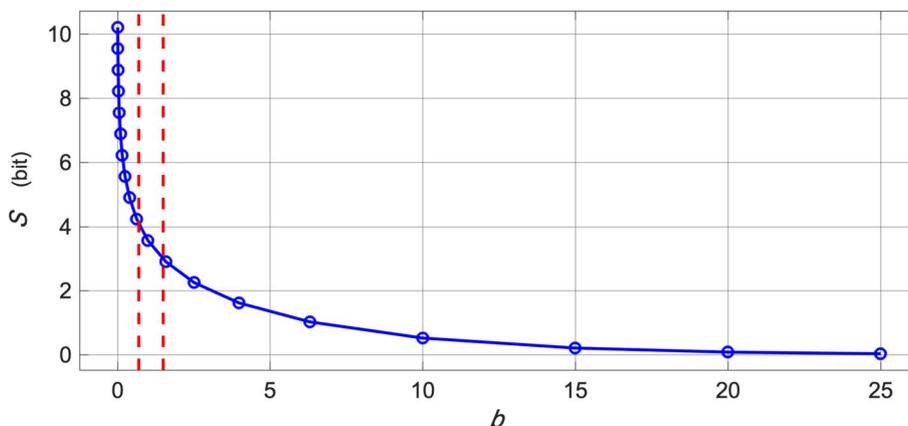
for $x = -\Delta M\beta$. Hence, the total entropy of an exponential distribution with parameter β expressed in bits is:

$$S = \beta\Delta M \frac{e^{-\beta\Delta M}}{1 - e^{-\beta\Delta M}} \log_2 e - \log_2 (1 - e^{-\beta\Delta M}). \quad (19)$$

Equation (19) is a closed, analytic expression for the information entropy of an exponential distribution with parameter β and class width ΔM . For a magnitude distribution, since $\Delta M = 0.1$ can be considered to be a set, constant value, (19) can be considered a direct relation between *S* and β (or $b = \beta \log_{10} e$). Although β has been used in the derivation of (19), results will be expressed in terms of *b*, because it is a more familiar parameter and its global average value, a good reference, is conveniently very close to 1.0 (e.g., El-Isa and Eaton 2014).

The direct, closed, relationship (19) between the *b*-value and the magnitude entropy is shown in Fig. 3. Figure 3

Fig. 3 Relationship between *b* and *S* (blue line), the dashed lines indicate the range of observed *b* values



also shows the range of entropies for reasonable b -values: from $S = 2.98$ bit for $b = 1.5$ to $S = 4.08$ bit for $b = 0.7$; a range of -1.1 bit for a b range of 0.8. This range has been chosen to illustrate the results because, although b -values in the $0.3 \leq b \leq 2.5$ range have been reported (El-Isa and Eaton 2014), for estimates based on magnitudes scaling as M_w (Hanks and Kanamori 1979; Kanamori 1983) $b = 0.7$ is an adequate lower limit for global b -values (Frohlich and Davis 1993) and an upper limit of $b = 1.5$ has proposed on physical grounds by Olsson (1999). Relationship (19) does not imply any maximum limit for b but tends asymptotically to zero as b grows.

Figure 3 shows that S increases as b decreases and, since as mentioned in the introduction b is inversely related to the state of stress in the medium, entropy appears to be directly related to said state of stress; indeed, since low b corresponds to probabilities being less concentrated around m_c , the significant probabilities are spread over a larger magnitude range, so the medium can be considered as being less ordered which means higher entropy.

Numerical entropies over a finite magnitude range

Now the results of (19) will be checked against numerical results from (4) to make sure that our entropy vs. b relationship is valid, and to see how results for finite K differ from those for $K \rightarrow \infty$. Although the G–R relation does not contemplate an upper limit for M , there are physical limits to how large a magnitude can be, so it is important to consider how results from a finite magnitude range differ from those of an infinite one. It is also important to consider the role of large magnitudes in the entropy determinations.

The very interesting problem of a maximum possible magnitude has been widely addressed (e.g., Beirlant et al., 2019; Chinnery and North 1975; Kijko 2004; Kijko and Singh 2011; Smith 1976; Sornette 2009) and manners of dealing with modified G–R distributions or using other distributions have been proposed (e.g., Cornell and Vanmarke, 1969; Cosentino et al. 1977; Holschneider et al. 2011; Lomnitz-Adler and Lomnitz 1979; Main and Burton 1984; Main 1996). The problem of a maximum magnitude is outside the scope of this paper, but it will be seen that the effects of very large magnitudes on entropy estimates are quite low and the possible existence, or not, of very large earthquakes does not affect the results shown here.

To corroborate the results of (19), the entropy of the magnitude distribution will be computed by evaluating exactly from (13) the probabilities for rounded magnitudes in a finite magnitude range, and using these probabilities to evaluate (4). The $2.0 \leq M_i \leq 9.0$ range has been chosen to illustrate the probabilities, because $M 2.0$ is not an

uncommon M_c and because $M 9.0$ is sufficiently rare as to be a practical upper limit because magnitudes much larger than 9.0 (including infinite ones) are not realistic.

Next, these exact theoretical probabilities will be used to calculate each term:

$$s_i \equiv s(M_i) = -P_i \log_2 P_i, \quad (20)$$

in the sum (4) and finally $S = \sum_{i=1}^K s_i$ will be computed and compared with the analytic total entropy values.

Figure 4 shows in (A) the theoretical probability mass distribution for three representative b -values; (B) shows the $s(M_i)$ corresponding to the probabilities shown in (A), and (C) shows the entropies computed using (4). Strictly speaking, the entropies correspond to the (larger) markers at the end of each curve, but the cumulative s leading to the total entropies is also shown, to illustrate its different behaviors for different b -values. The dotted horizontal lines in (C) correspond to the analytical entropies.

For the smallest magnitudes s is largest for the higher b , but about one magnitude unit above M_c the roles are reversed and the entropies for smaller b -values grow faster and soon the entropy for the smallest b is the largest of all. All

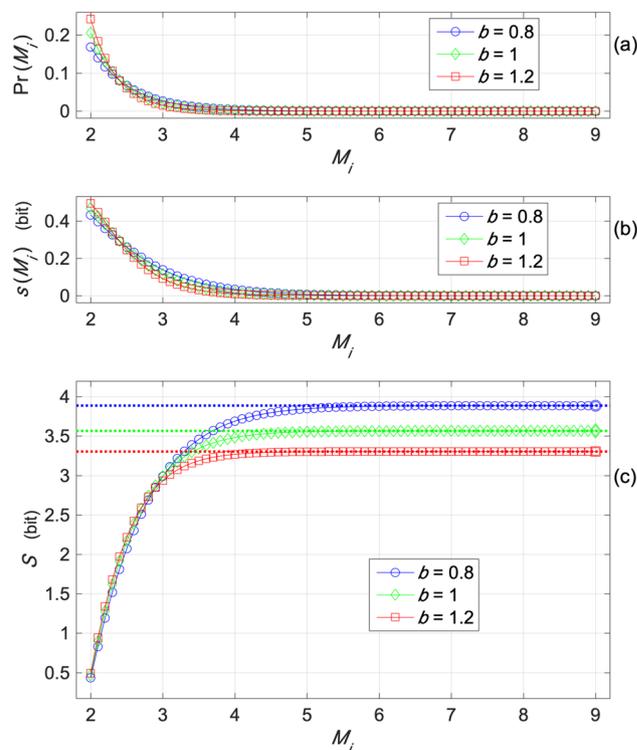


Fig. 4 Exact numerical probabilities **a**, corresponding information scores **b** and entropies **c**, for three representative b -values and a finite magnitude range. Panel **c** shows the numerical entropy values as large symbols over the largest magnitude, and the analytical entropies as dotted lines; also shown are the cumulative s values

entropies tend asymptotically to their theoretical values, with the largest b -values approaching it earlier. The magnitudes that make more difference are those in the $3.0 \leq M \leq 5.0$ range.

As shown in Fig. 4 (A), large magnitudes have very small probabilities which are close to the left end of the shaded area in Fig. 1 and contribute very little to the total entropy, as shown in (B) and (C). Hence, the presence of magnitudes above 6.5 or 7.0 is not necessary for obtaining good, approximate estimates of S .

The numerical values for the total entropies differ from the analytic ones by only 4.2×10^{-5} for $b = 0.8$, 2.0×10^{-6} for $b = 1.0$, and 9.0×10^{-8} for $b = 1.2$, differences too small to be of practical concern. As would be expected from the properties of the exponential distribution, shifting the magnitude range while conserving the same width, to $1.5 \leq M_i \leq 8.5$, say, results in exactly the same entropy estimates.

Estimates do change if the range is enlarged, for example, considering the $1.5 \leq M_i \leq 9.0$ range (five classes wider) reduces the differences between numerical and analytical to 1.7×10^{-5} for $b = 0.8$, 6.7×10^{-7} for $b = 1.0$, and 2.4×10^{-8} for $b = 1.2$, because of the contributions from the extra five terms in (15).

For reference, the entropy of a uniform distribution with K classes is:

$$S_U = - \sum_{i=1}^K \frac{1}{K} \log_2 \frac{1}{K} = \log_2 K, \quad (21)$$

so for the example, with range $2.0 \leq M_i \leq 9.0$ and $K = 71$, the entropy of the uniform distribution, i.e., the largest possible entropy, would be $S_U = 6.15$ bit, some 2.26 bit larger than the entropy for $b = 0.8$.

The total entropies are distinctly larger for the smaller b -values, which means that measuring entropies can be a good method for identifying regions of low or large b , that is, of large or low stress.

Numerical entropy from samples

Next, it will be seen how entropy measured from samples compares to the entropy computed from exact probabilities, and how it depends on the sample size; the samples will be synthetics from random simulations, for the same magnitude range and the three representative b -values used above.

For each b -value, N exponentially distributed random magnitudes are generated as:

$$m = m_c - \ln(1 - r * \rho) / \beta, \quad (22)$$

where r is a uniformly distributed pseudo-random number in the zero to one range, and

$$\rho = 1 - e^{\beta(m_c - m)} \quad (23)$$

maps this range onto the range that results in probabilities $m_c \leq m \leq m_x$.

With these magnitudes, a histogram with classes ΔM wide, corresponding to the rounded magnitudes, is constructed and the number of events in each class $n(M_i)$ is counted. The probabilities are estimated as:

$$P_i = n(M_i) / N \quad (24)$$

(c.f. Feng and Luo 2009) and used in (20) to calculate the s_i values and thence S .

Figure 5 shows simulations for three b -values, each having $N = 5000$ magnitudes, a reasonably good-sized sample. The magnitude histograms $n(M_i)$ are shown in (A), and the contributions $s(M_i)$ are shown in (B); the cumulatives for s and the entropies are shown in (C), together with the theoretical entropies.

A comparison of panels (C) of Figs. 4 and 5 shows very good agreement between entropies from theoretical and simulated magnitudes, both converging nicely to the analytic entropies from (19).

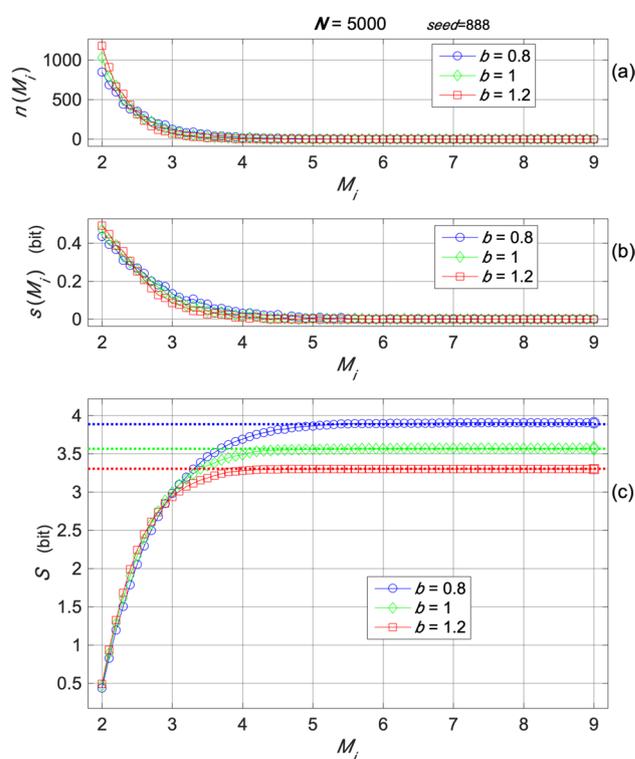


Fig. 5 Numerical probabilities from a synthetic sample of 5,000 magnitudes **a**, corresponding information scores **b** and entropies **c**, for three representative b -values and a finite magnitude range. Panel **c** shows the numerical entropy values as large symbols over the largest magnitude, and the analytical entropies as dotted lines; also shown are the cumulative s values

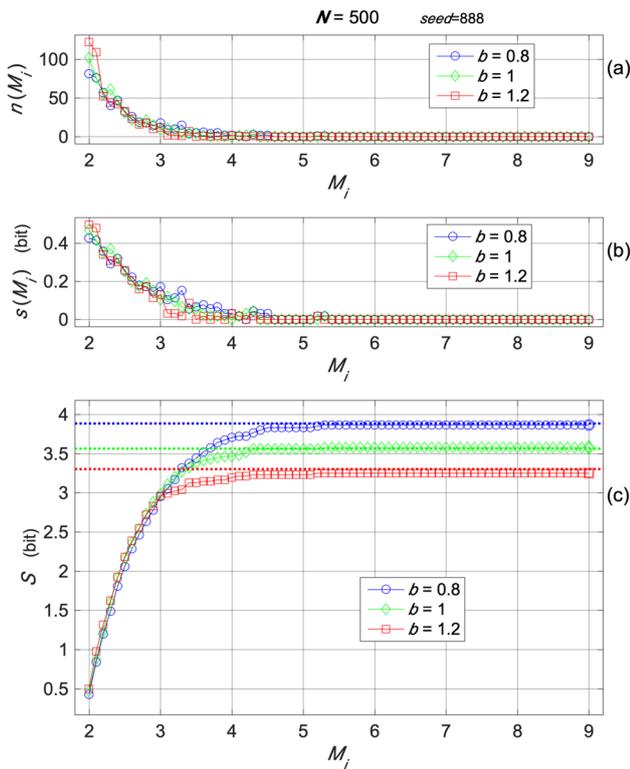


Fig. 6 Numerical probabilities from a synthetic sample of 500 magnitudes **A**, corresponding information scores **B** and entropies **C**, for three representative b -values and a finite magnitude range. Panel **C** shows the numerical entropy values as large symbols over the largest magnitude, and the analytical entropies as dotted lines; also shown are the cumulative s values

The simulations shown in Fig. 6 are like those of Fig. 5, but for a much smaller sample of $N = 500$ magnitudes. The histograms in the (A) and (B) panels show clear differences from the respective graphs in Fig. 5; differences are less apparent between panels (C), but there is a noticeable difference for the entropy corresponding to the largest $b = 1.2$, which is well below the analytic entropy.

Monte Carlo simulations and sample size

Monte Carlo simulations are used to characterize how numerical entropies depend on sample size, each simulation consisting of $N_r = 5000$ realizations, like those shown in the previous section, of magnitude samples of different sizes, from $N = 250$ to $N = 5000$. The means and standard deviations of the N_r entropies calculated for each combination of b and N are shown in Fig. 7.

Figure 7 shows the mean calculated S as a thick line with a particular color and symbol for each b -value and shows the mean plus/minus one standard deviation as thin lines and the true analytical value as a dotted line in the corresponding color. In order to interpret correctly the information in

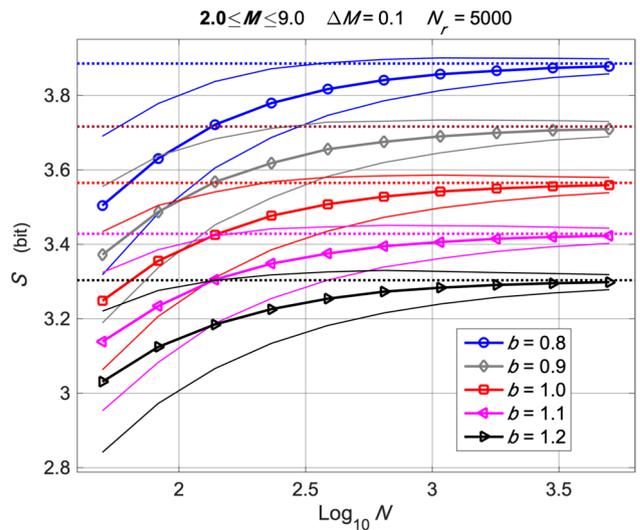


Fig. 7 Monte Carlo analysis of entropies S determined from synthetic samples for different sample sizes N . Thick lines with different colors and symbols, corresponding to representative b -values are the means of 5,000 realizations for each combination of b and N . The thin lines show the means plus/minus one standard deviation, and the horizontal dotted lines indicate the analytical entropies

the standard deviations it is necessary to determine how the entropy values are distributed, and Fig. 8 shows an example of these distributions for $b = 1.0$ and $N = 5000$, which tells us that the values can be considered to be normally distributed around the mean.

Figure 7 shows that the entropy estimated from samples smaller than ~ 200 will almost certainly be undervalued, particularly for low b . Entropies corresponding to b -values differing by as much as 0.1 cannot be distinguished with 0.7 confidence for samples smaller than about 350 for low b

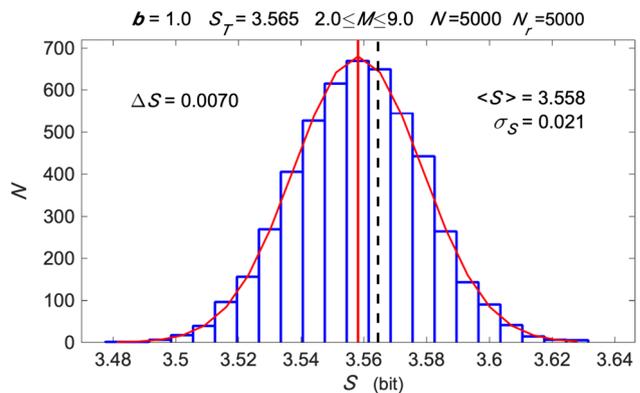


Fig. 8 Histogram of $N_r = 5,000$ Monte Carlo entropy determinations for $b = 1.0$ and magnitudes in the $2.0 \leq M \leq 9.0$ range (blue line); the vertical red line shows the mean value, and the vertical dashed line is the analytical S value. The thin line is a normal distribution for the observed standard deviation σ_S multiplied by N_r

and about 550 for high b , and distinguishing them with 0.95 confidence requires $-1,500$ and $-3,00$ samples, respectively.

For samples -2000 to -2500 , mean values underestimate the analytical entropy by -0.01 bit, and for samples of 5000 the underestimations go from 0.0067 bit for $b = 0.8$ to 0.0051 bit for $b = 1.2$, with standard deviations -0.02 bit. For the larger samples, the means tend to the analytical entropies very slowly, and including larger magnitudes does not help very much because their number is very small and, as shown in Figs. 1, 5, and 6, their contribution to the total entropy is almost insignificant.

Standard deviations diminish slowly, and even for large samples $-5,000$ the standard deviation corresponding to $b = 0.8$, $\sigma_S = 0.0201$, is -0.005 of the mean value $\bar{S} = 3.8778$, while for $b = 1.2$, $\sigma_S = 0.0202$ is -0.0061 of the mean value $\bar{S} = 3.2978$. These normalized standard deviations are smaller than the corresponding ones for b -values estimated by the Aki–Utsu method for the same synthetic samples used to evaluate the entropies.

Figure 7 shows that, although entropies evaluated over a finite magnitude range should be smaller than the analytical ones, the entropies measured from samples could be overvalued and thus be slightly larger.

Measured entropies and b -values for single trials

It has been discussed how entropies are measured from data, and Figure 7 shows how the measurements can be expected to agree with the real values, but in practice the real values are not known nor are there thousands of realizations; usually the data correspond to a single realization and there is no way of knowing how well it conforms to the behavior of the means shown above.

Since there is an explicit relation between S and b , it would seem that their measures would be redundant but this is not exactly the case because they are measured in different ways. b -value measurements (2) depend only on \bar{M} , while entropy estimations depend on the values of all entropy scores s_i .

In order to illustrate how single realizations agree with, or differ from, the means of many realizations and from the true values, let us look at four examples of single sample realizations, and see how each single realization depends on sample size. All realizations share exactly the same parameters and differ only in the number used as seed for the pseudo-random number generator. Each realization was a set of $N_T = 5,000$ magnitudes, and we obtained estimates of S , using (24) and (4), and b , using (2), for subsets of $N = 500, 600, 700, \dots, 5,000$, and from each b , we calculated the entropy using (19).

The examples are shown in Fig. 9, where panels (C) plot the histograms of the total N_T magnitudes to show that the synthetic magnitudes are indeed exponentially distributed. Panels (B) show the estimated b -values and, for reference, the true b -value, while panels (A) show the estimated entropies as blue circles, the analytic entropy corresponding to the true b , and show as asterisks the entropies computed from the estimated b -values.

As mentioned above, the realizations in Fig. 9 differ only in the random number seed and illustrate how a realization corresponding to some set of real data can vary randomly while being a product of a given conditions on a given seismic system. The two upper examples show “expected” behaviors, with values varying considerably for short samples and gradually converging to a value close to the true one, albeit one (upper left) from above and the other (upper right) from below. The example at lower left does converge but does not reach the true value, and the example at lower right does not converge to the true value at all. It should be said that most realizations behave more like the good examples, so that many different seeds were tried before the ugly example at lower right was obtained.

All the examples show that for small data sets the measured entropies and those estimated from the b -values differ very much for small samples, but run almost parallel for large samples. Entropies from b estimates are larger than measured ones, but that is to be expected because of the finite magnitude range. Thus, it is proposed that, although related and calculated from the same data, entropy and b measurements are not just scaled versions of each other, because they are calculated in different ways that are sensitive to different kinds of errors, and when both measurements are correct they should agree within the limitations. Hence, the differences between directly estimated entropies and those estimates from b -values can help us determine when samples are adequate and results are trustworthy.

As an example of how the entropy and Aki–Utsu b -value estimates are not necessarily equivalent, consider the contribution of a very large magnitude. Let a data set have $N - 1$ elements, and let the entropy determined from the sample be:

$$S_{[N-1]} = \sum_{i=1}^{N-1} p_i \log_2 p_i, \quad (25)$$

and the b -value be:

$$b_{[N-1]} = \frac{\ln 10}{\bar{M}_{[N-1]} - m_c}, \quad (26)$$

where $\bar{M}_{[N-1]} = \frac{1}{N-1} \sum_{i=1}^{N-1} M_i$. Now, let the next magnitude M_N be large enough so that it stands alone in a class, then,

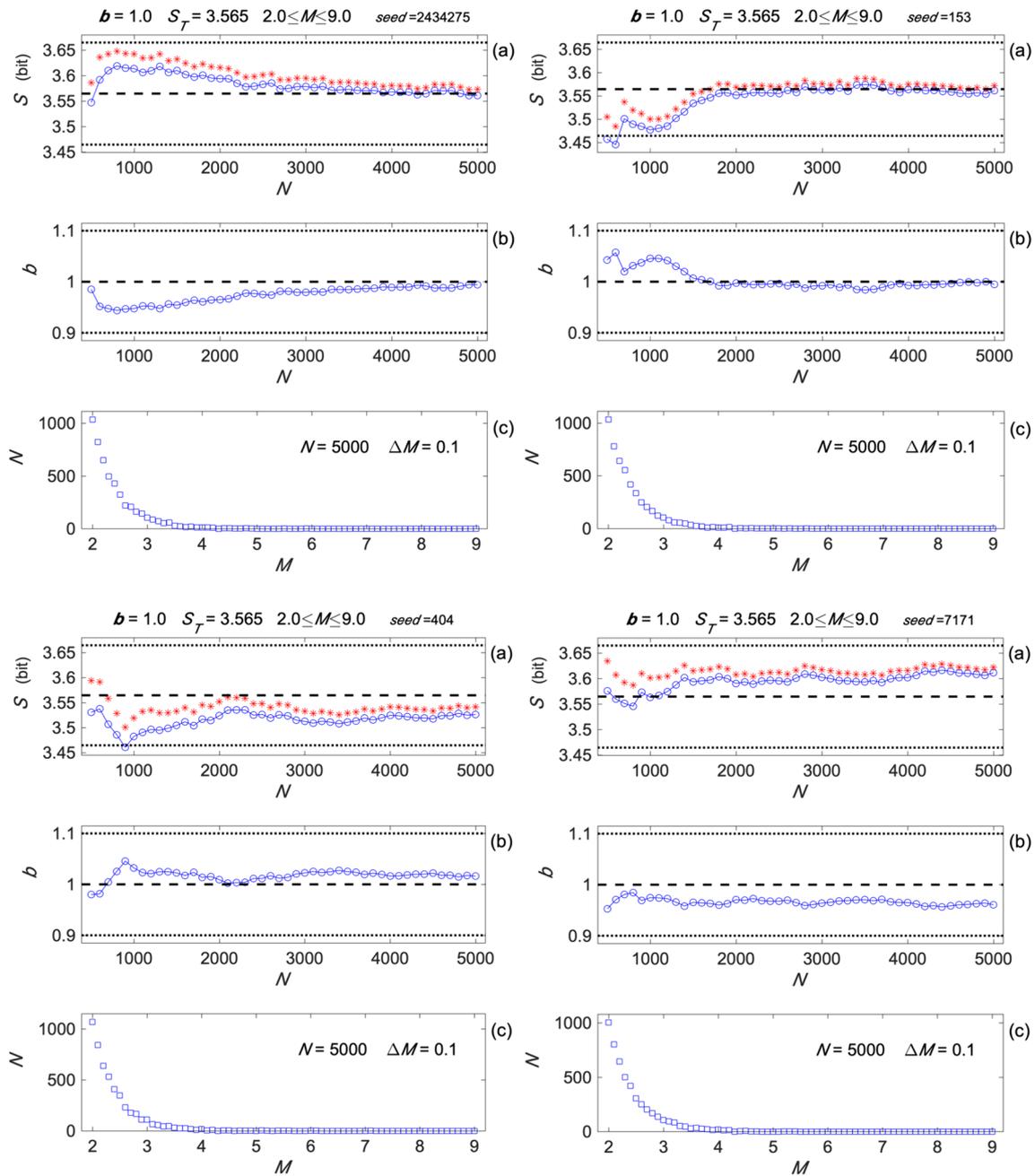


Fig. 9 Four examples of entropy and b -value determinations from single realizations of $N_T = 5,000$ synthetic exponentially distributed magnitudes, taken N elements at a time. The **c** panels show the magnitude histograms for the total N_T data, the **b** panels show the b -values estimated using (2) with \bar{M} determined from N data

(blue circles) and the true b -value (dashed red line). The **a** panels show as blue circles the entropies determined for each N data, and as asterisks the entropies estimated from the measured b -values in **b**, using (19); the analytical entropy corresponding to the true b is shown as a dashed red line

because there is only one event in the class, its probability will be $1/N$, so

$$S_{[N]} = \frac{N-1}{N} S_{[N-1]} + \frac{N-1}{N} \log_2 \frac{N-1}{N} + \frac{1}{N} \log_2 \frac{1}{N} \quad (27)$$

and the change of entropy does not depend on the value of M_N , as long as it is large enough to stand alone in a class. On the other hand,

$$b_{[N]} = \frac{\ln 10}{\bar{M}_{[N-1]} - m_c + \frac{1}{N} (M_N - \bar{M}_{[N-1]})} \quad (28)$$

does depend on the actual value of M_N . Hence, unless N is very large, the effect of a large magnitude is different for entropies and for b -values.

Discussion and Conclusions

A closed analytical relationship between the b -value (or β) that characterizes the magnitude G–R distribution, or any other exponential distribution, and the information entropy of the distribution has been derived (19). The relationship was checked by means of the numerical evaluation of the entropy computed using the exact probabilities derived from the distribution.

Neither the G–R distribution nor the associated exponential distributions contemplate a maximum magnitude, it was possible to evaluate the effect of working with a finite magnitude range on the entropy, and it was found that, because very small probabilities contribute very little to the entropy, the difference between the analytical and the finite range entropies is quite small.

Next, the results of the relationship were compared with entropies estimated from synthetic sets of exponentially distributed random data, and very good agreement was found.

Using Monte Carlo simulations, the accuracy and precision for entropy evaluations as a function of sample size were explored. The evaluations were found to be distributed normally around their means, which allows setting familiar confidence limits to the power of discriminating between different values of the entropy.

Although b -values and entropies are formally related, their evaluations from the data are done by different methods and so are affected differently by different characteristics of the data, particularly for small data sets. Hence, it is proposed that entropy and G–R b -value measurements can be complementary and help to estimate when a sample is large enough for results to be reliable.

Acknowledgements Our sincere thanks to Editor Ramón Zúñiga and to two anonymous reviewers.

Declarations

Conflict of interest The author has not disclosed any conflict of interest.

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